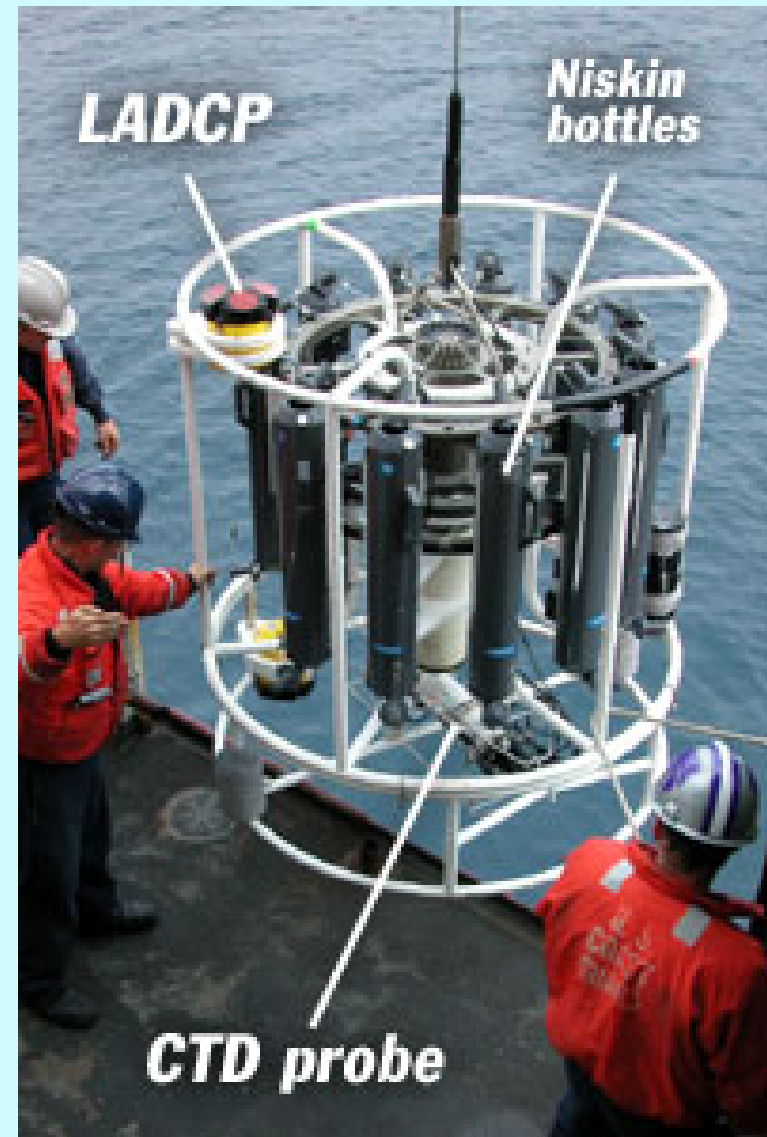
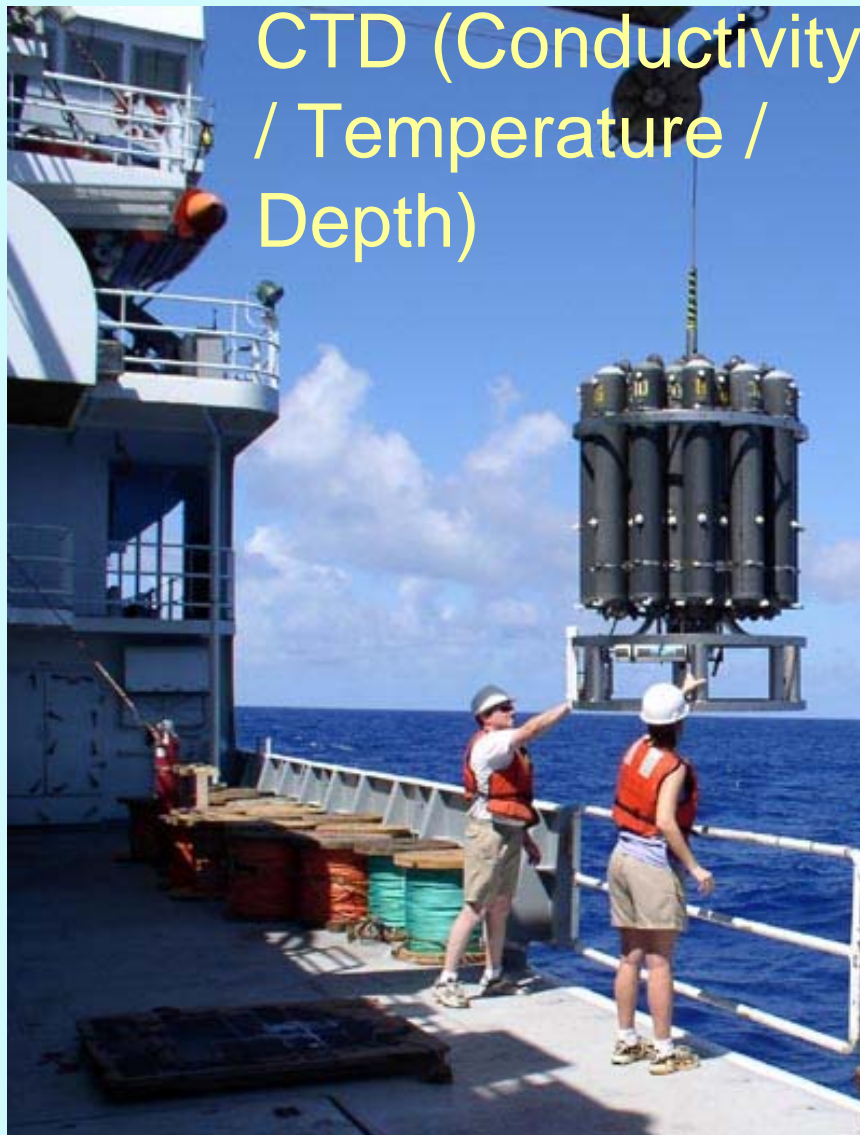
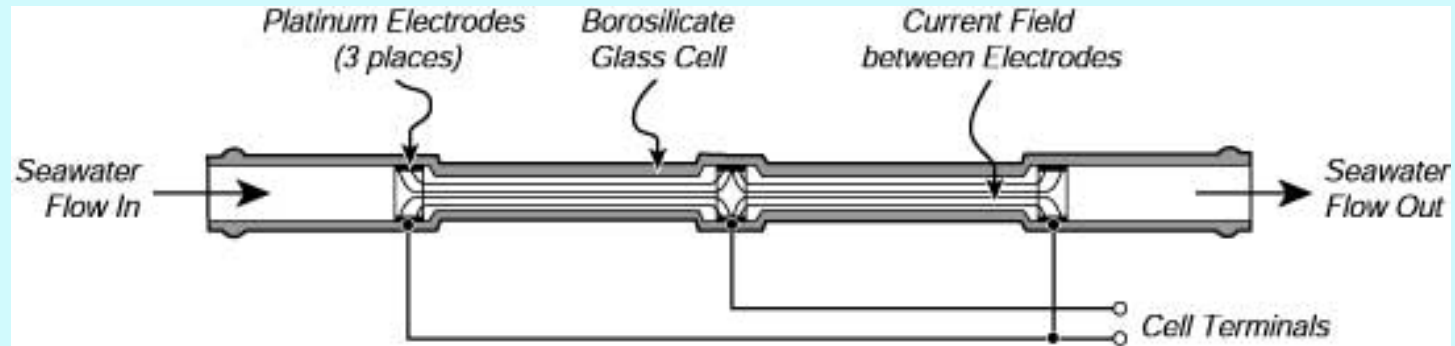


# AdOc 4060 / 5060 Seawater Salinity

2013 Spring  
Chris Jenkins



# Conductivity measurement



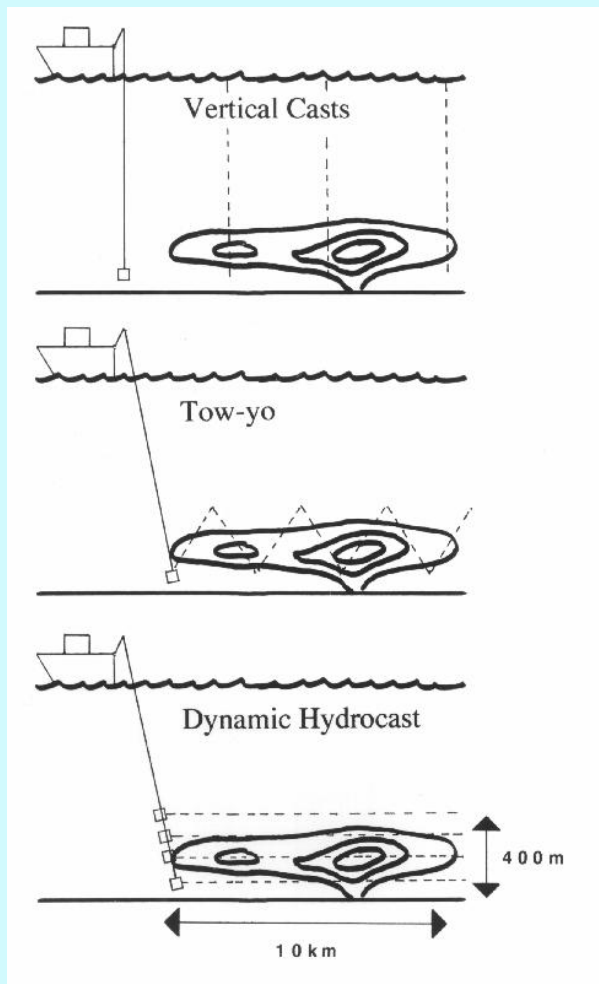
- Accuracy: 0.75 psu (practical salinity unit)
- Ocean water: 35 psu = 35,000 ppm (by weight) of salt in water

# Temperature measurement

- Thermistor (Pt)
- Range: -1 to 40 deg C
- Accuracy: +/- 0.1deg C



# Profiling



# Niskin bottles



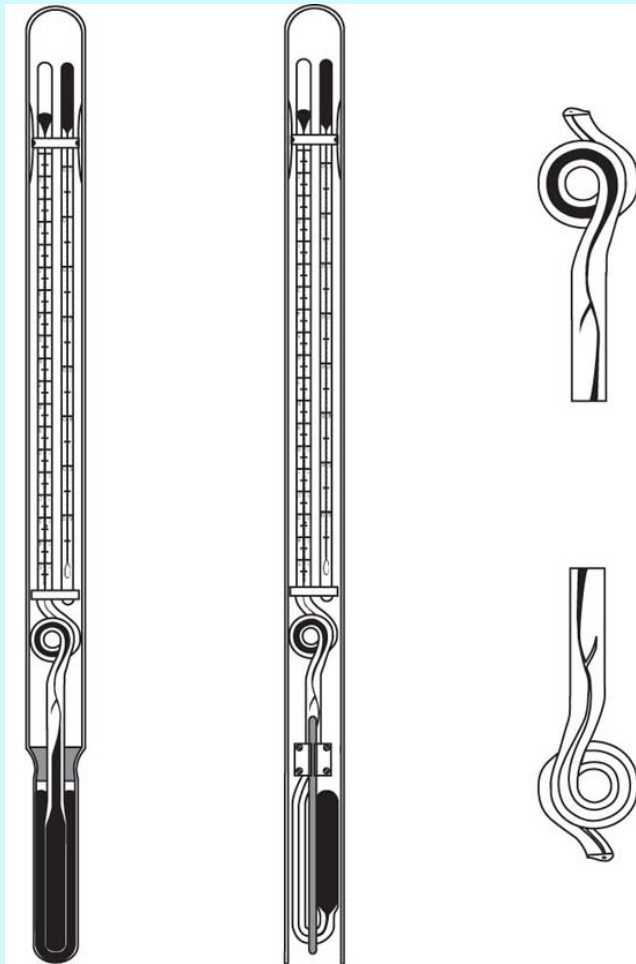
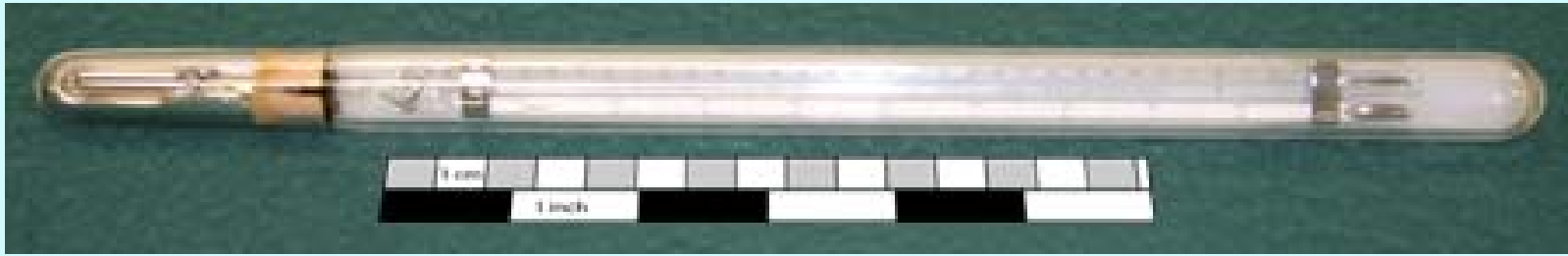
- Bottles can be closed when desired by researchers on the ship
- Water samples can be obtained to calibrate the CTD and to investigate properties not measured by the CTD such as dissolved nutrient content, carbon dioxide, chlorofluorocarbons, and others tracers.

# Nansen bottles



**Nansen water bottles  
before (I), during (II), and  
after (III) reversing.  
(From Dietrich et al. 1980)**

- Approximately 12 to 24 Nansen bottles are attached in series at predetermined intervals along a cable
- Deployed with both ends open
- Closes as the bottles are reversed
- Used in conjunction with inverting or reversing thermometers



## Inverting thermometers

- Has a constriction in the capillary that allows it to retain its reading upon being inverted
- Deployed in pairs, one protected and the other unprotected from high pressures. Combination of the two gives temperature and pressure at the depth where reversal took place

# Salinity

- The total amount of solid material dissolved in water.
- Concentration, often described as parts per thousand
- Average ocean salinity is 35 parts per thousand (or about 3.5%)
  
- By evaporation of seawater (inaccurate)
- By chlorinity method (via titration)  
$$S = 1.80655 \ C/$$
- By electrical conductivity  
conductivity is proportional to salinity

# Main constituents of seawater (ions)

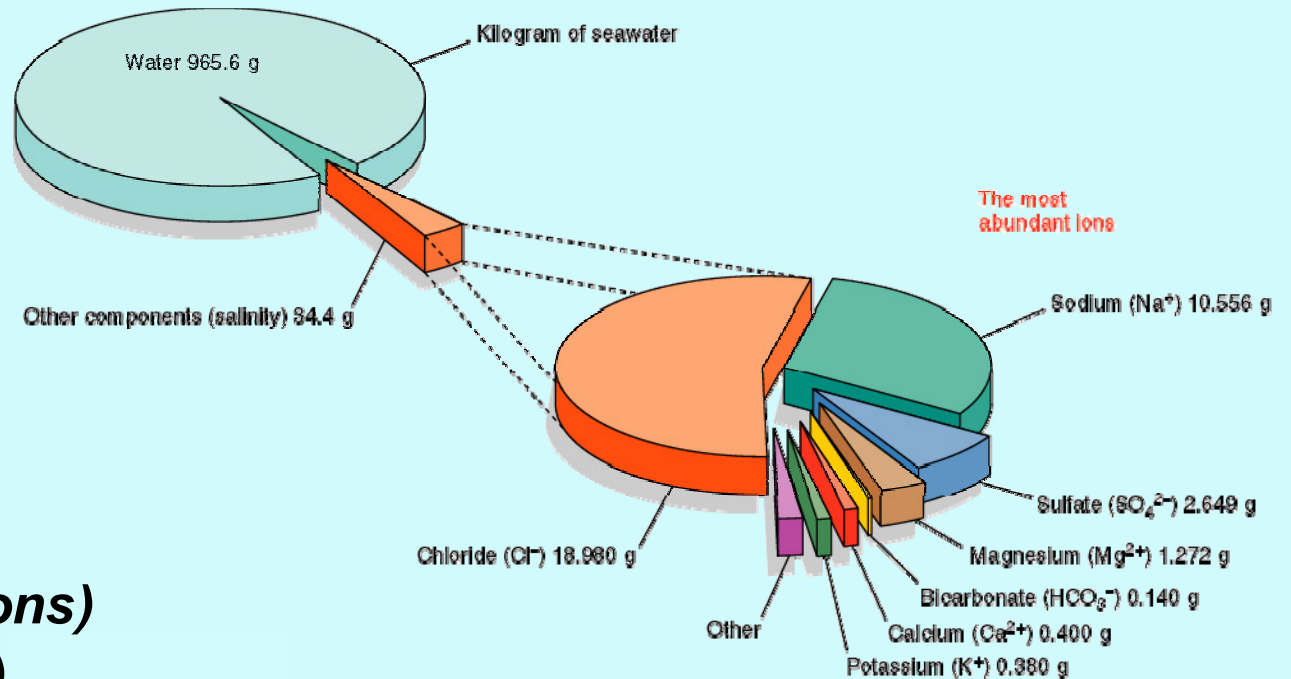
Main constituents of seawater (> 1ppm):

## ***anions (-ve ions)***

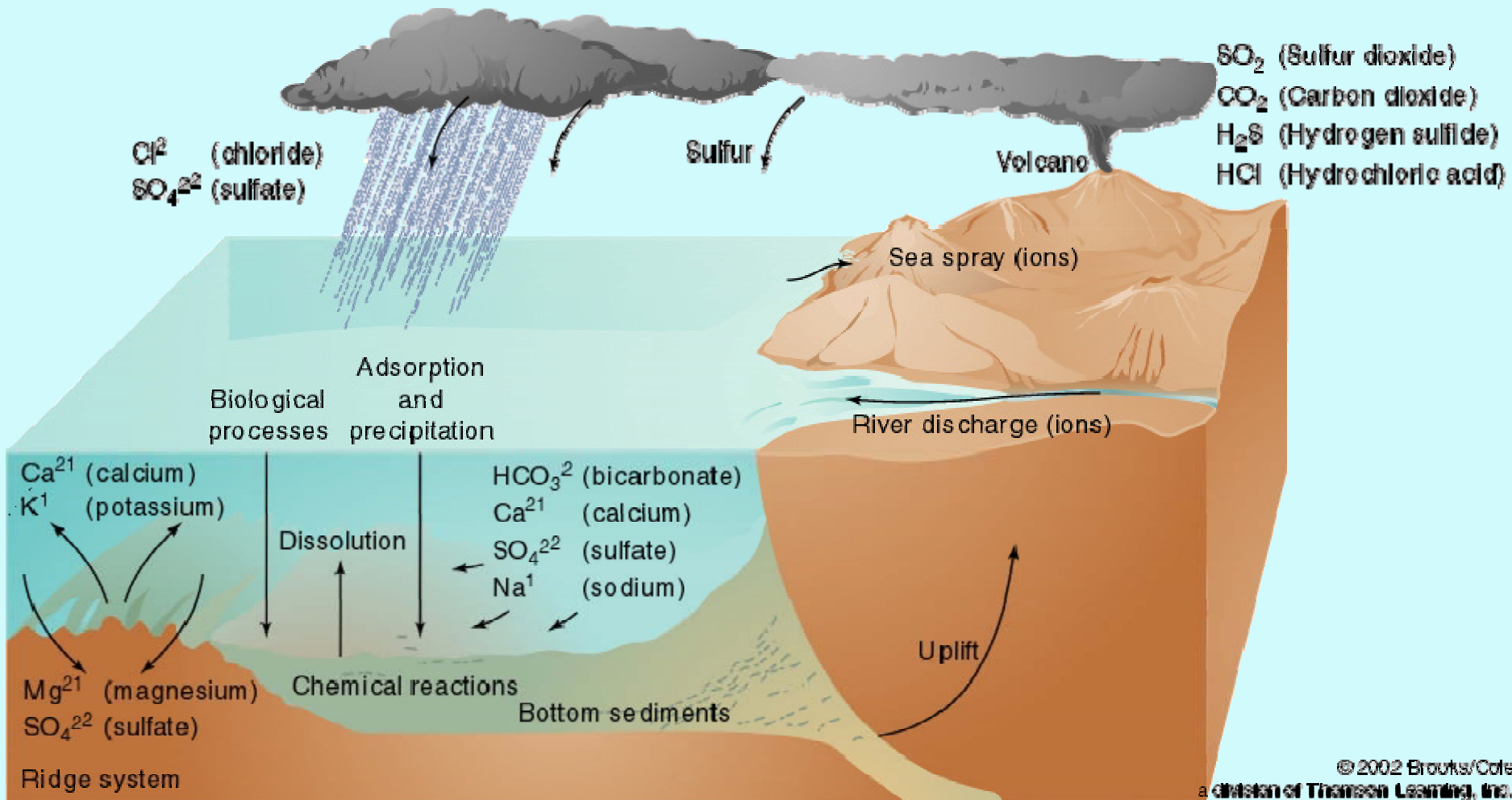
- Chloride
- Sulphate ( $\text{SO}_4^{2-}$ )
- Bicarbonate
- Bromide,
- Borate & Fluoride

## ***cations (positive ions)***

- Sodium ( $\text{Na}^+$ , cation)
- Magnesium ( $\text{Mg}^{2+}$ )
- Calcium
- Potassium
- Strontium







**Table 7.3 Approximate Residence Times for Constituents of Seawater**

<b>Constituent</b>	<b>Residence Time (years)</b>
Chloride ( $\text{Cl}^-$ )	100,000,000
Sodium ( $\text{Na}^+$ )	68,000,000
Magnesium ( $\text{Mg}^{2+}$ )	13,000,000
Potassium ( $\text{K}^+$ )	12,000,000
Sulfate ( $\text{SO}_4^{2-}$ )	11,000,000
Calcium ( $\text{Ca}^{2+}$ )	1,000,000
Carbonate ( $\text{CO}_3^{2-}$ )	110,000
Silicon (Si)	20,000
Water ( $\text{H}_2\text{O}$ )	4,100
Manganese (Mn)	1,300
Aluminum (Al)	600
Iron (Fe)	200

*Sources:* Data from Broecker and Peng, 1982; Bruland, 1983; Riley and Skirrow, 1975.

# Forchhammer, Dittmar and the law of constant proportions

- Regardless of salinity, the major ions are found in the same proportions throughout the world ocean (chemical equilibrium)
- This means that if you measure one, you have them all.
- Oceanographic insights can be found in the exception to this rule.

# Example: Law of constant proportions

Suppose you are given a sea water sample and you are told that it has 12.1 ppt of Sodium.

What is the overall salinity of this sample ?

Table 3.1 (p. 29) shows that standard seawater has:

$$\begin{aligned}\text{Sodium (ppt) / Salinity (ppt)} &= 10.556 / 34.482 \\ &= 0.306\end{aligned}$$

Our sample must have this same ratio, so we have:

$$12.1 \text{ (ppt) / Salinity} = 0.306$$

With a little algebra we find:            Salinity = 39.54 (ppt)

## CONCLUSIONS

Using Newly generated data, a fit has been made giving the following algorithm for the calculation of salinity from data of the form:

$$R = \frac{C(S, T, P)}{C(35, 15, 0)}$$

$T$  in °C (IPTS '68),  $P$  in decibars.

$$R_T = \frac{R}{R_{PT}}; R_P = 1 + \frac{P \times (A_1 + A_2 P + A_3 P^2)}{1 + B_1 T + B_2 T^2 + B_3 R + B_4 R T}$$

$$r_T = c_0 + c_1 T + c_2 T^2 + c_3 T^3 + c_4 T^4$$

$$A_1 = 2.070 \times 10^{-5} \quad B_1 = 3.426 \times 10^{-2}$$

$$A_2 = -6.370 \times 10^{-10} \quad B_2 = 4.464 \times 10^{-4}$$

$$A_3 = 3.989 \times 10^{-15} \quad B_3 = 4.215 \times 10^{-1}$$

$$B_4 = -3.107 \times 10^{-3}$$

$$c_0 = 6.766097 \times 10^{-1}$$

$$c_1 = 2.00564 \times 10^{-2}$$

$$c_2 = 1.104259 \times 10^{-4}$$

$$c_3 = -6.9698 \times 10^{-7}$$

$$c_4 = 1.0031 \times 10^{-9}$$

$$S = \sum_{j=0}^5 a_j R_T^{j/2} + \frac{(T-15)}{1+k(T-15)} \sum_{j=0}^5 b_j R_T^{j/2}$$

$$a_0 = 0.0080 \quad b_0 = 0.0005 \quad k = 0.0162.$$

$$a_1 = -0.1692 \quad b_1 = -0.0056$$

$$a_2 = 25.3851 \quad b_2 = -0.0066$$

$$a_3 = 14.0941 \quad b_3 = -0.0375$$

$$a_4 = -7.0261 \quad b_4 = 0.0636$$

$$a_5 = 2.7081 \quad b_5 = -0.0144$$

# Practical Salinity Scale PSS 1978

<http://www.iode.org/oceanportal/detail.php?id=3540>

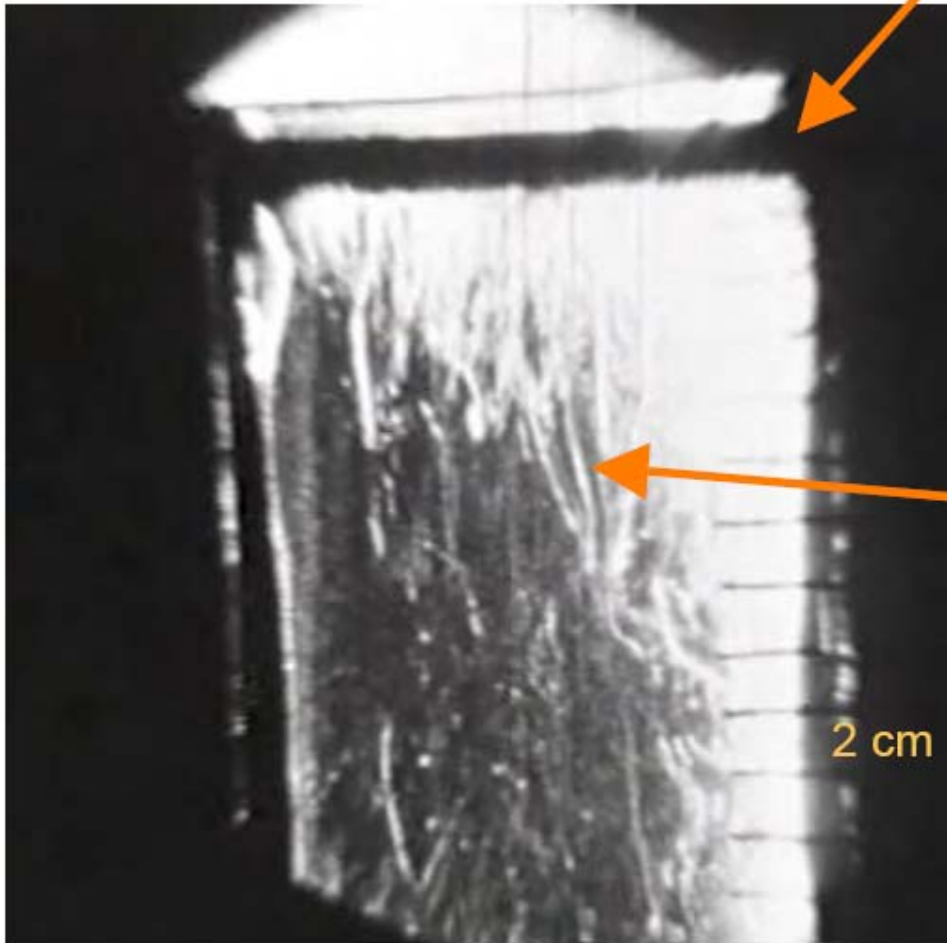
# Salinity variations

Location/type	Salinity
Normal open ocean	33-38 ‰
Baltic Sea	10 ‰ (brackish)
Red Sea	42 ‰ (hypersaline)
Great Salt Lake	280 ‰
Dead Sea	330 ‰
Tap water	0.8 ‰ or less
Premium bottled water	0.3‰
Human tears	9 ‰

# Rejected Brine and Brine Plumes

*Shadowgraph*

growing sea ice



Movie of Wakatsuchi,  
1974 / 1983

brine plumes

2 cm

# Processes affecting seawater salinity

- Processes that decrease seawater salinity:
  - Precipitation
  - Runoff
  - Icebergs melting
  - Sea ice melting
- Processes that increase seawater salinity:
  - Sea ice forming
  - Evaporation

Precipitation and evaporation are most important overall.



# Surface salinity variation

- Pattern of surface salinity:
  - Lowest in high latitudes
  - Highest in the tropics
  - Dips at the Equator
- Surface processes help explain pattern

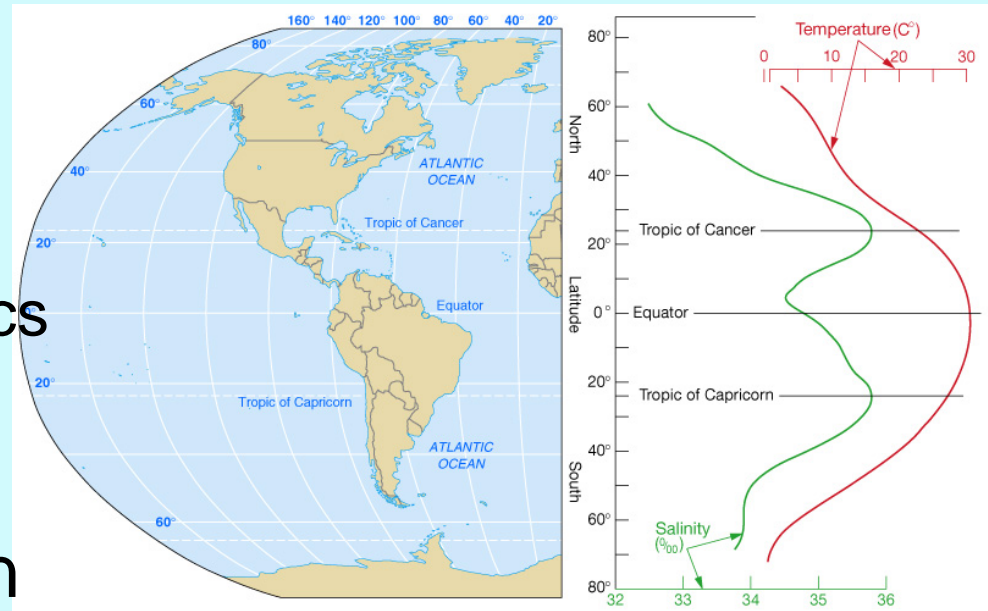
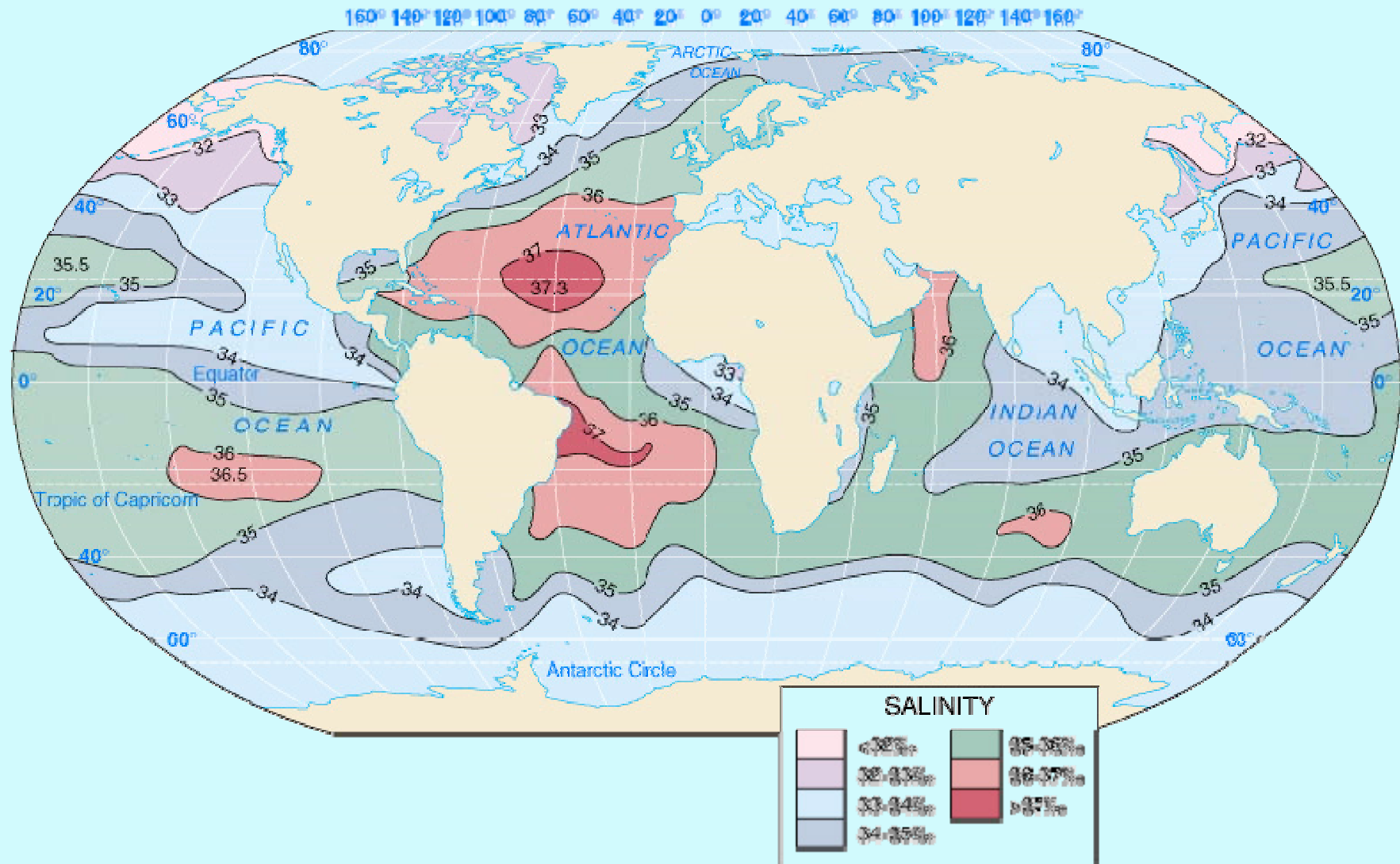


Figure 5-20

# Surface salinity variation

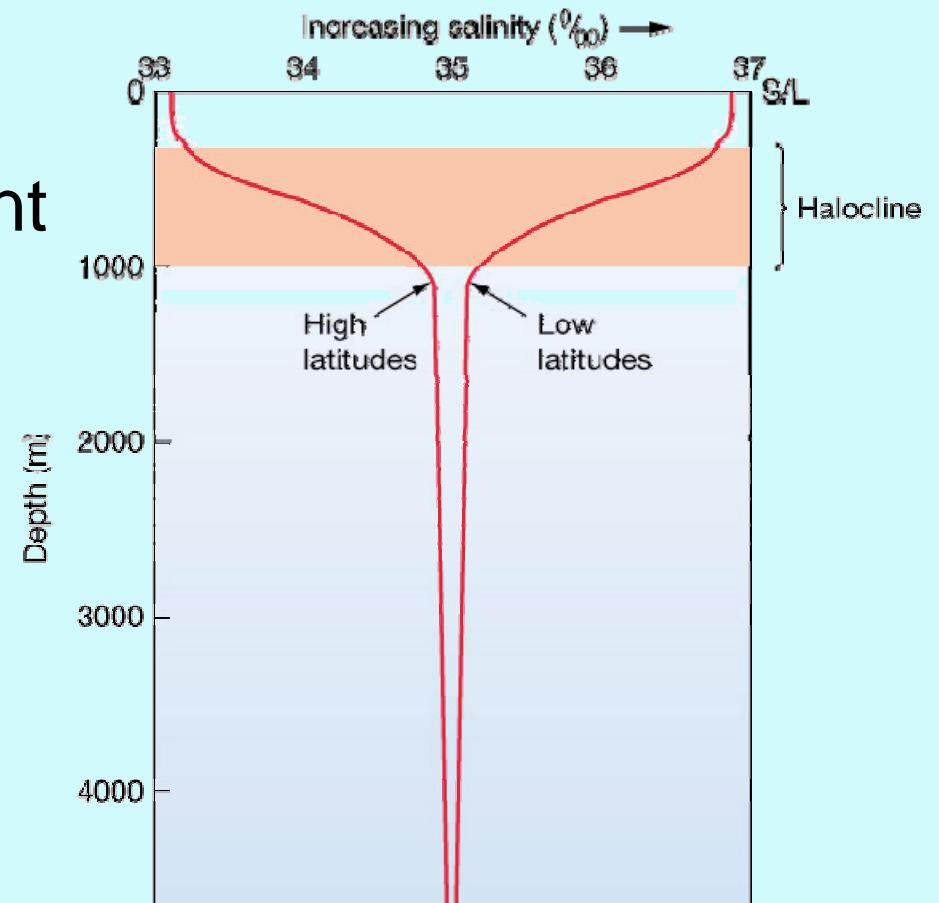
- High latitudes have low surface salinity
  - High precipitation and runoff
  - Low evaporation
- Tropics have high surface salinity
  - High evaporation
  - Low precipitation
- Equator has a dip in surface salinity
  - High precipitation partially offsets high evaporation

# Global surface salinity

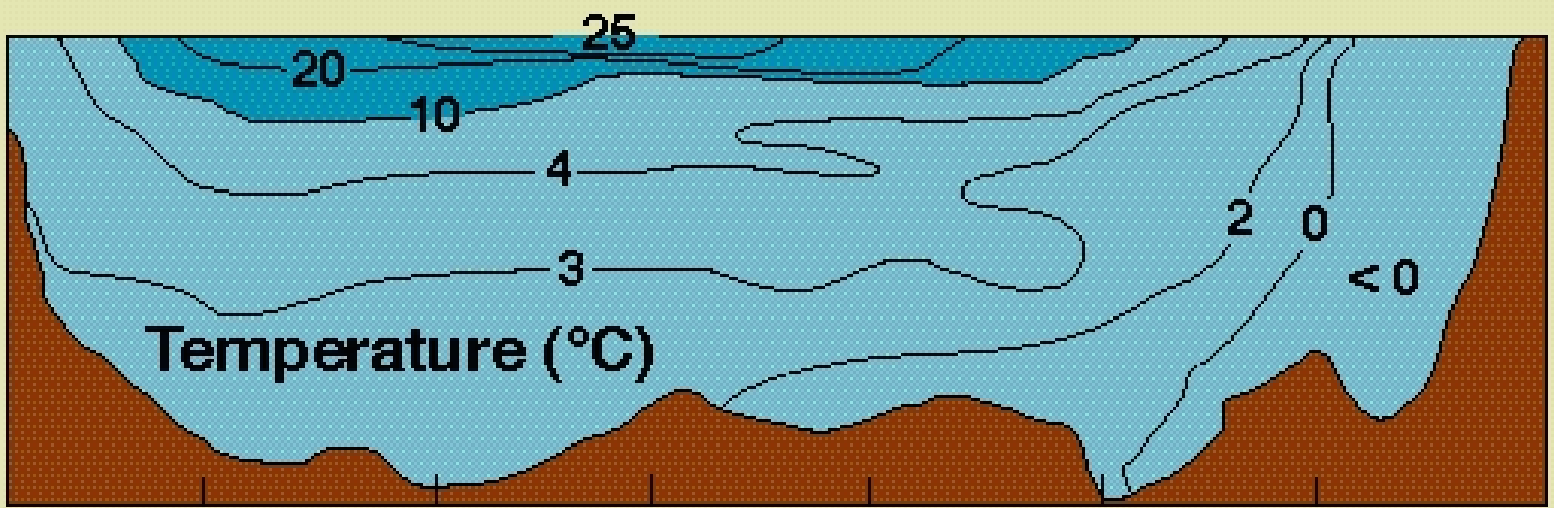
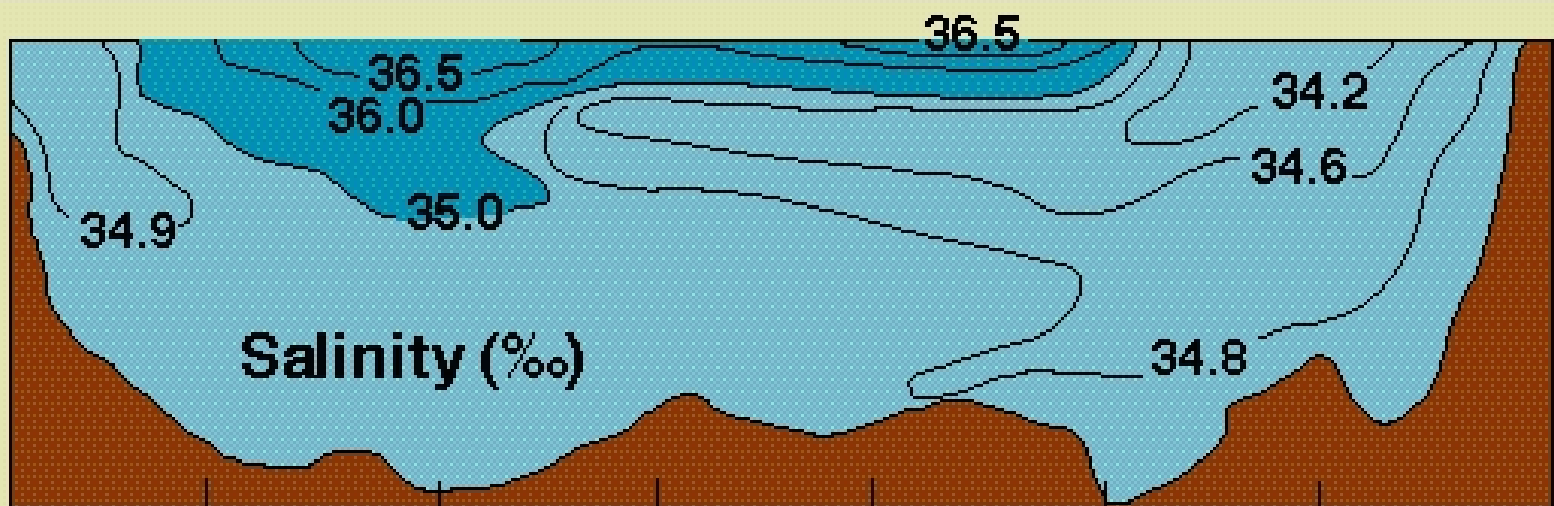


# Salinity variation with depth

- Curves for high and low latitudes begin at different surface salinities
- **Halocline** = layer of rapidly changing salinity
- At depth, salinity is uniform



# Atlantic Ocean Salinity/Temperature Structure



60°N 40° 20° 0° 20° 40° 60° 80°S

## Summary

- SW salinity is on average 35 ppt by weight
- 11 major ions make up 99.9% of the dissolved constituents
- Constituents differ greatly from crustal rock due to differences in their solubilities
- Salinity varies with depth and latitude but the constituent ratios remain virtually constant: evaporation and precipitation change the total salinity but not the composition
- Minor departures are a result of biological processes affecting  $\text{Ca}^{++}$  and  $\text{HCO}_3^-$
- Major departures occur only locally (anoxic conditions, hydrothermal vents)
- Surface salinities are largest in the tropical and subtropical latitudes where evaporation is greater than precipitation; and smallest in higher latitudes where precipitation and runoff is greater than evaporation
- Electrical conductivity is the modern means of measurement

# Salinity clock ?

